



NCAR

Remote Sensing of Water Vapor Using Microwave Radiometer

Presentation to the 3rd SoMEX/TiMREX
Science Workshop

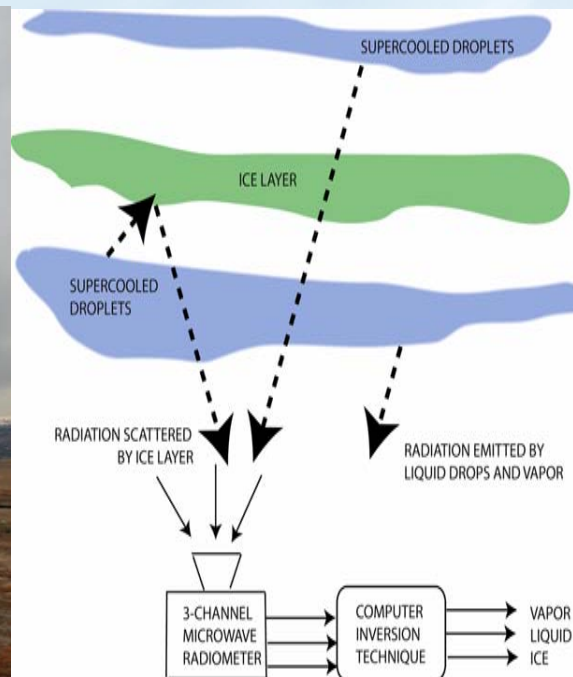
Taipei, Taiwan
November 4, 2010

Vivek
Earth Observing Laboratory
National Center for Atmospheric Research
Boulder, Colorado, USA

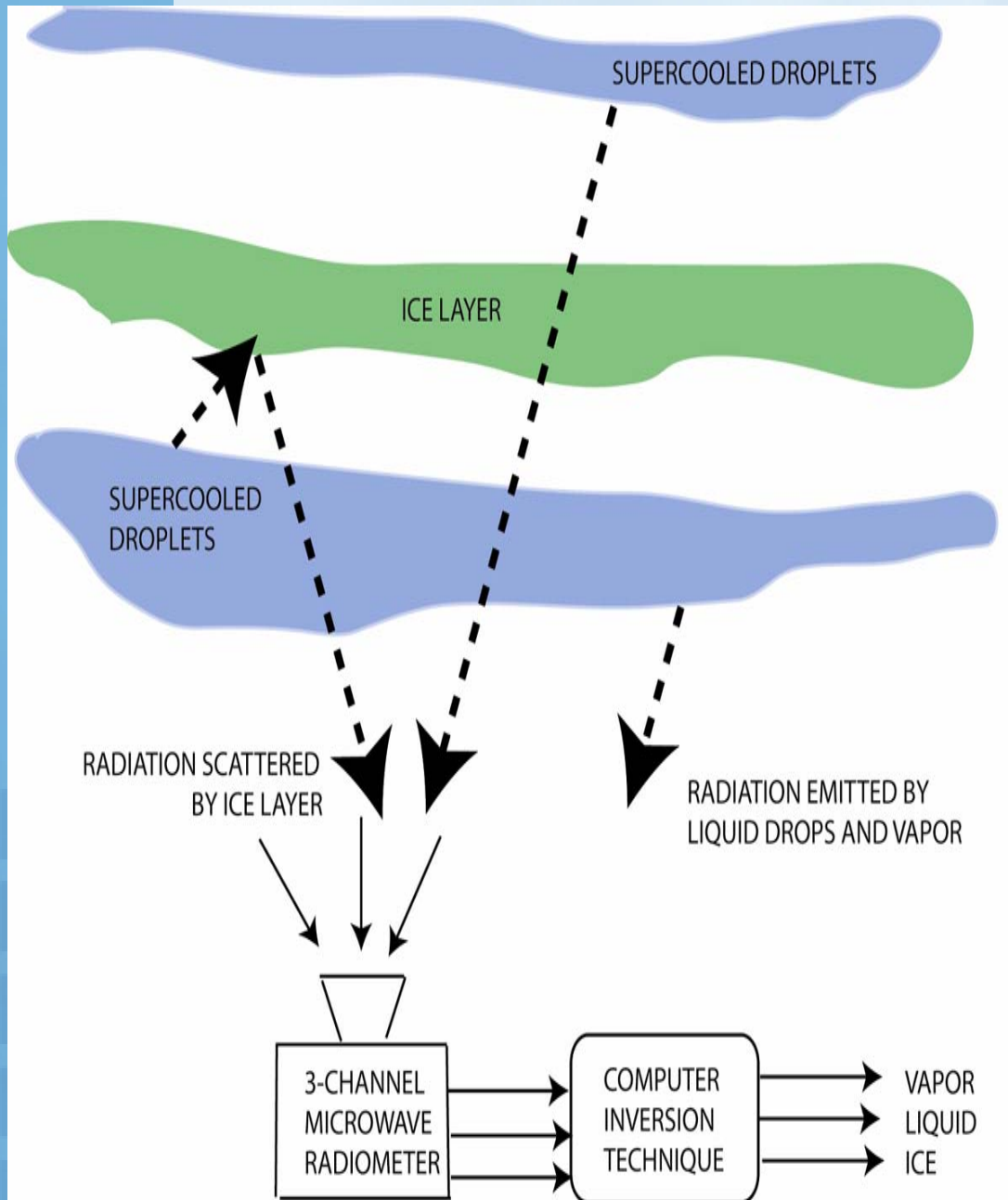
SwaroopSahoo and Steve Reising
Colorado State University
Fort Collins, Colorado, USA

Instruments

- Ground-based radar: (i) total phase or delay and (ii) absorption using a dual-wavelength radar
- Ground-based multi-wavelength microwave radiometer
- Satellite-borne microwave radiometer



Radiometer Technique



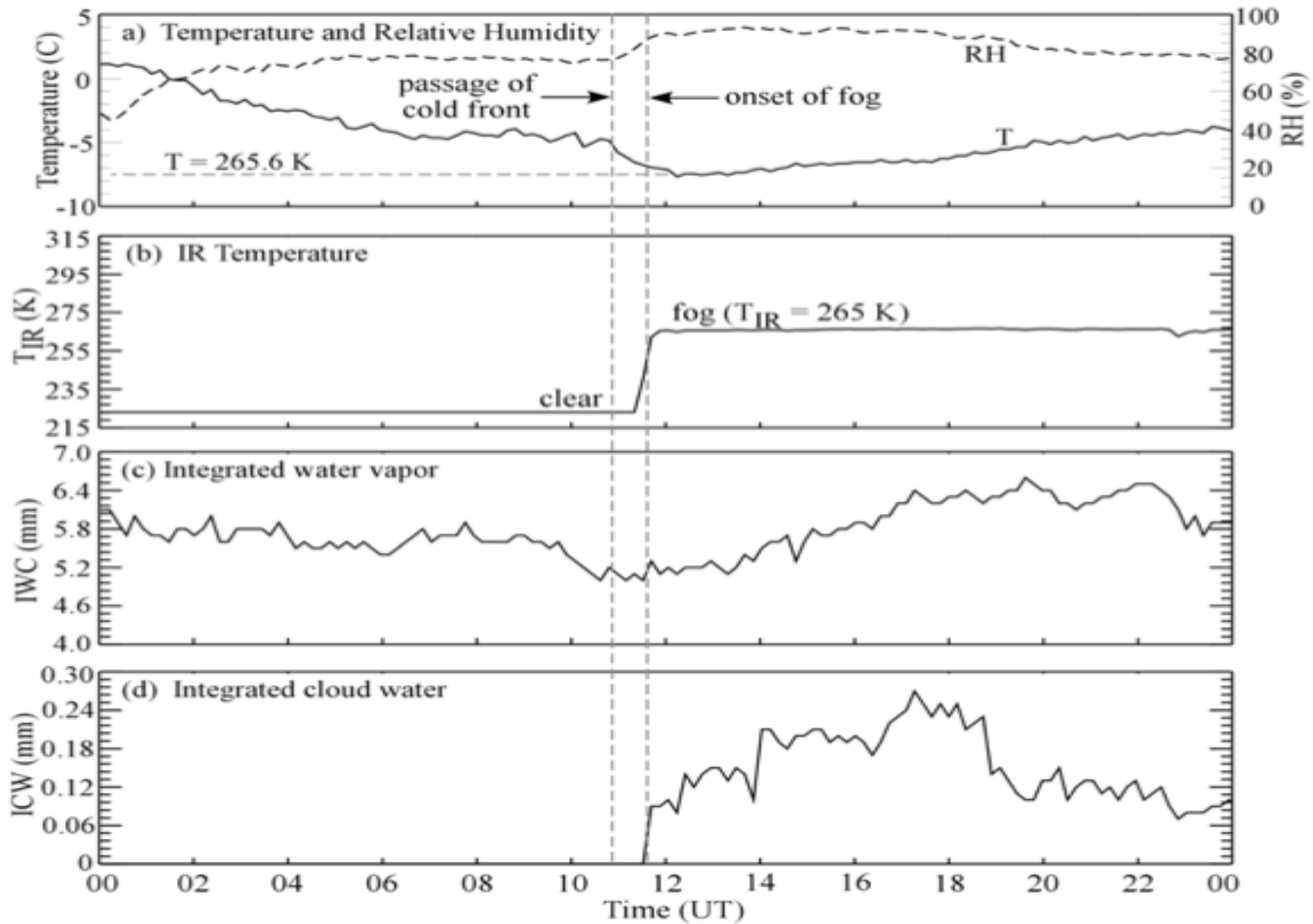
- Radiometers measure brightness temperatures T_b , that are converted into optical depths, τ .

$$\tau = \log \left(\frac{T_m - T_{\text{sky}}}{T_m - T_b} \right)$$

- Optical depths are linearly related to LWP and VWP
- k_l and k_v are path averaged coefficients.
- τ_d is the 'dry' optical depth
- Two wavelengths, two equations, two unknowns - retrieve LWP and VWP.

$$\tau = k_l \text{LWP} + k_v \text{VWP} + \tau_d$$

Times series measurements of temperature, integrated vapor and liquid



Spatial and Temporal Resolutions



(a) Radiosondes:

- Synoptic scale; twice daily. Thermodynamic vertical profiles for weather diagnostics and prediction
- Don't resolve microscale and mesoscale features of minutes to hours and 1-10 km scale

(b) Satellite:

Crude vertical resolution within boundary layer

(c) Radiometer: Continuous observations to fill temporal gaps between radiosondes

(i) Microwave: Measurements during both cloudy and clear air

(ii) Infrared: Biased in cloudy condition

Vertical resolution declines in proportion to height above ground level



NCAR

Ground-based radiometer

Profiling Using a Microwave Radiometer

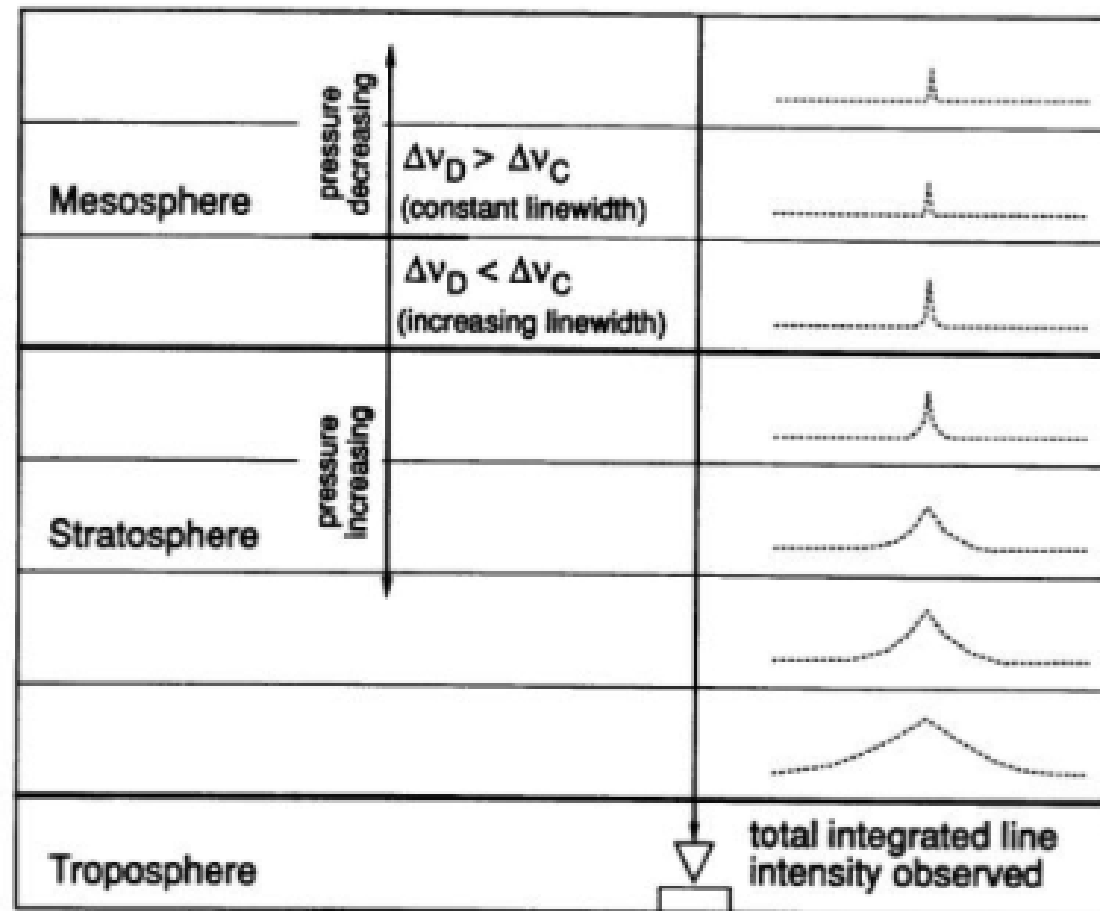
Frequency of the emitted radiation
Bohr's Equation

$$\nu_{mn} = \frac{E_m - E_n}{h}$$

Total internal energy of a molecule:

$$E_{int} = E_e + E_v + E_r$$

Two important rotational transitions: 22.235 and 183.31 GHz



Brightness Temperature and weighting function

T_B : Brightness temperature
 $W(f,z)$: Weighting function
 $g(z)$: Water vapor density

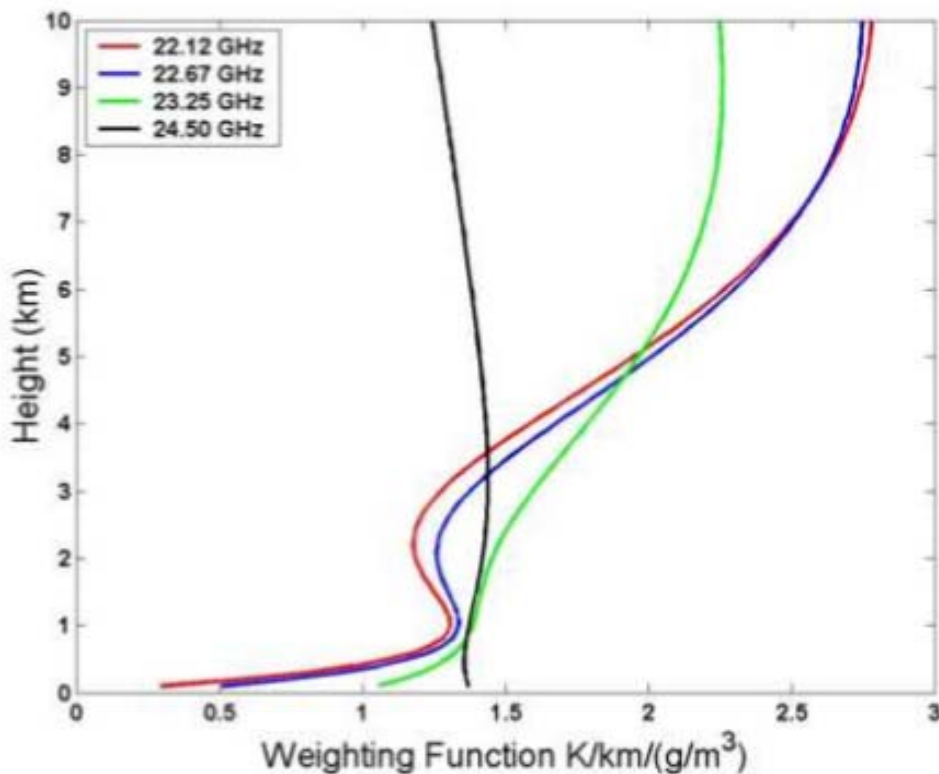
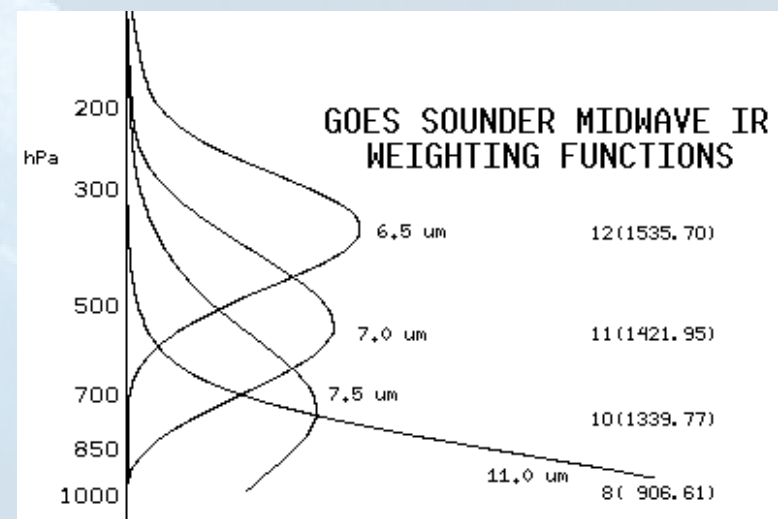


Figure 3.1: Weighting function at the four CMR-H frequencies



Corrections for finite beamwidth and mean radiating temperature



T_B : Corrected brightness temperature

Retrieval Method: 1-D VAR (Variational Assimilation Retrieval)



W : $m \times n$ matrix

m : # of measurements

n : # of altitudes at which the WV density is desired

$m < n$ i.e. # of measurements are smaller than # of unknowns

$\rho_{v,a}$: a priori water vapor profile

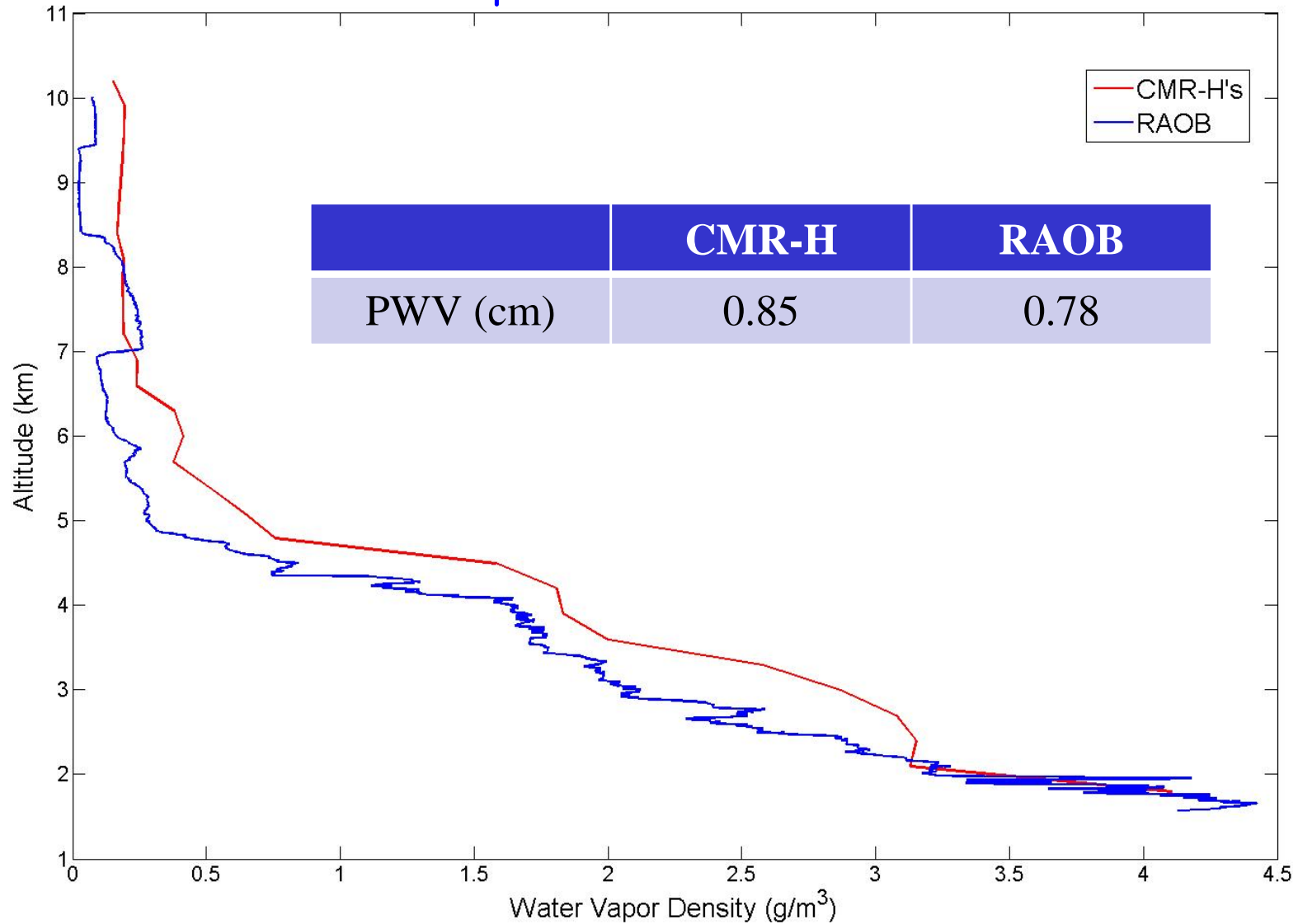
: Error covariance matrix of the a priori water vapor profile

: Error covariance matrix of the measurements

Comparison of RAOB and CMR-H Retrieved Water Vapor Density Profile on Oct 9, 2007



09 Oct 07 at 6 UT. A priori from Denver Station RAOB at 0 UT



Mathematical model for describing spatial resolution

Measurement:

.... (1)

Retrieved WV:

.... (2)

Retrieved WV = f(weighting function, retrieval coeff., measurement errors)

Brightness Temperature and weighting function

T_B : Brightness temperature
 $W(f,z)$: Weighting function
 $g(z)$: Water vapor density

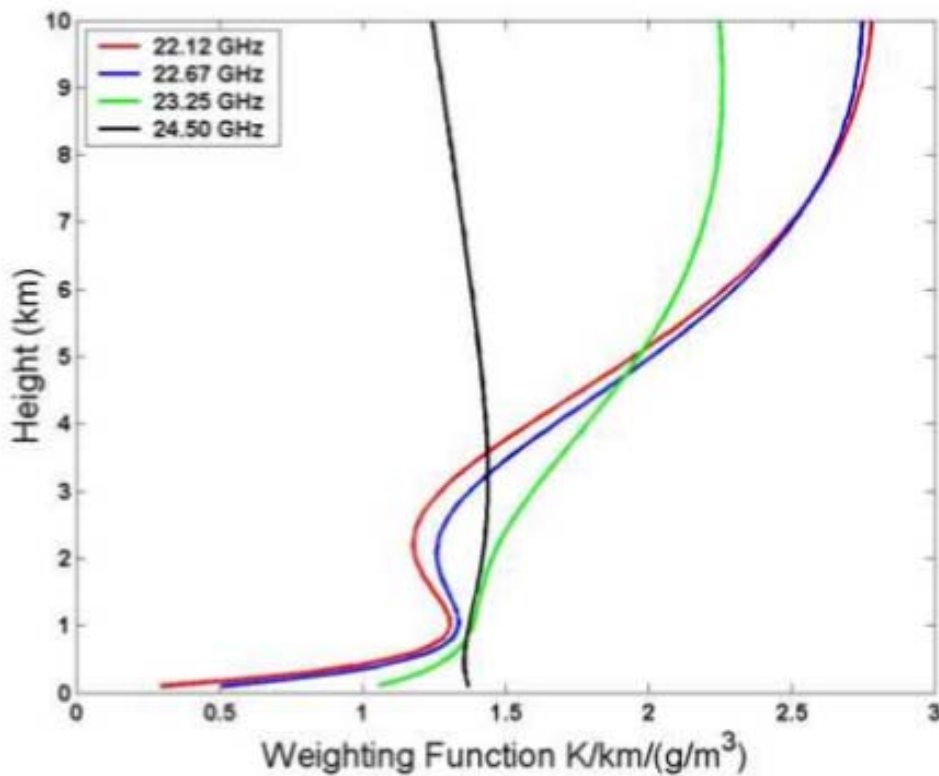
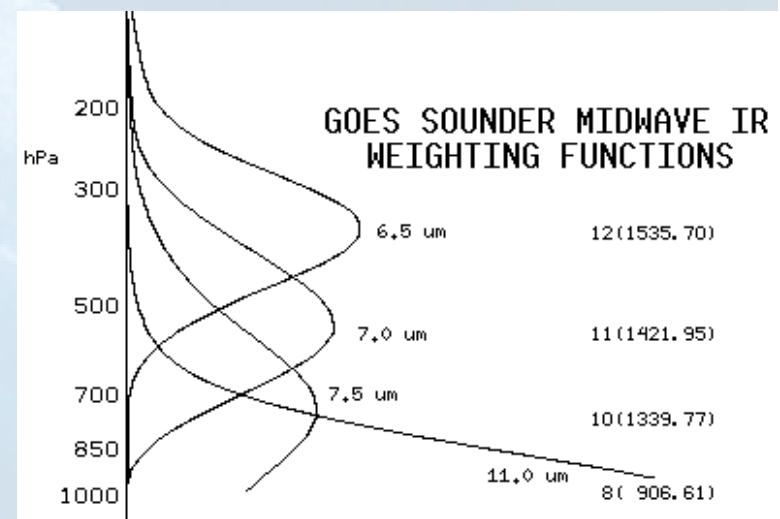


Figure 3.1: Weighting function at the four CMR-H frequencies



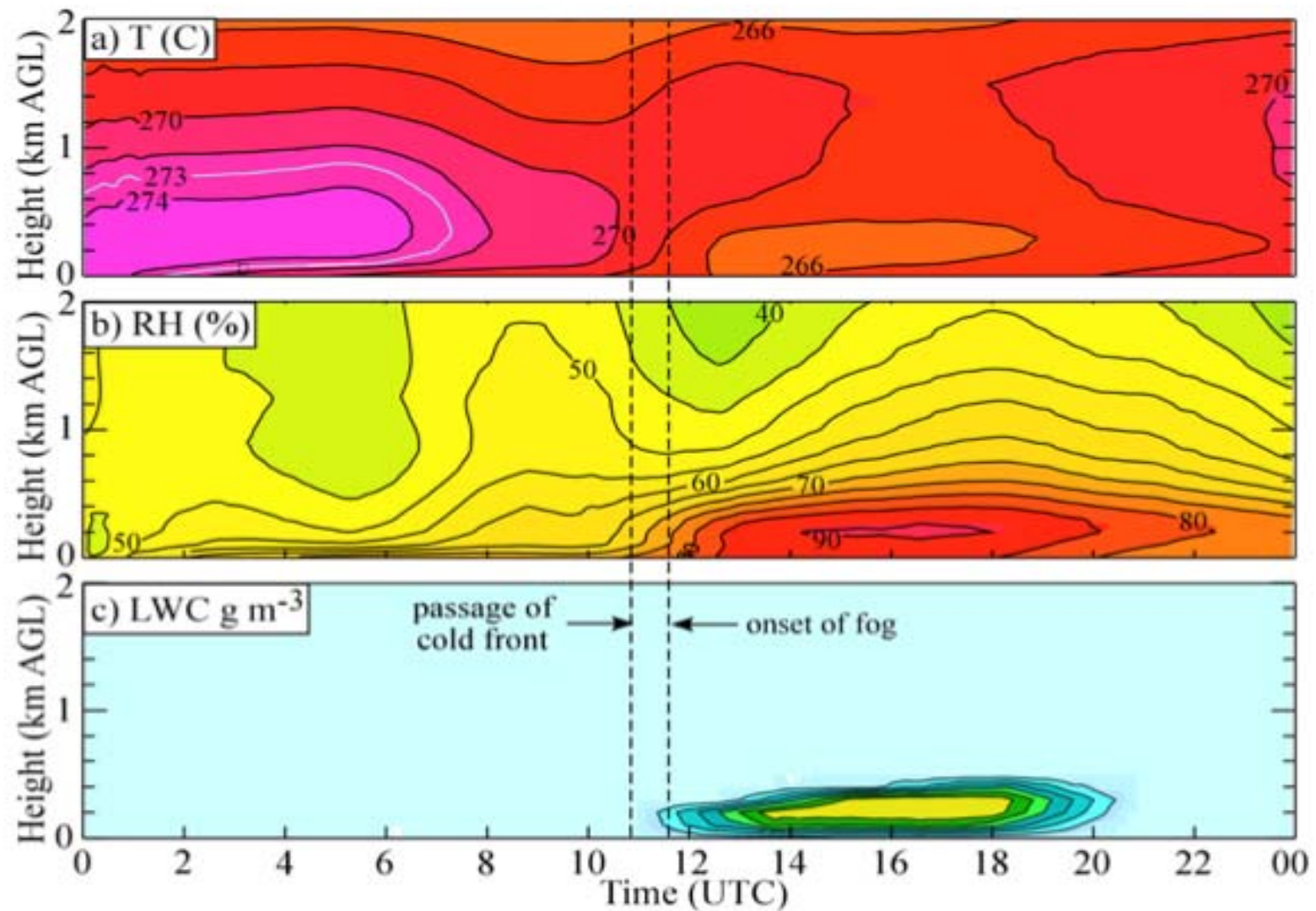
Improvement in resolution by adding scanning measurements to vertical pointing observations



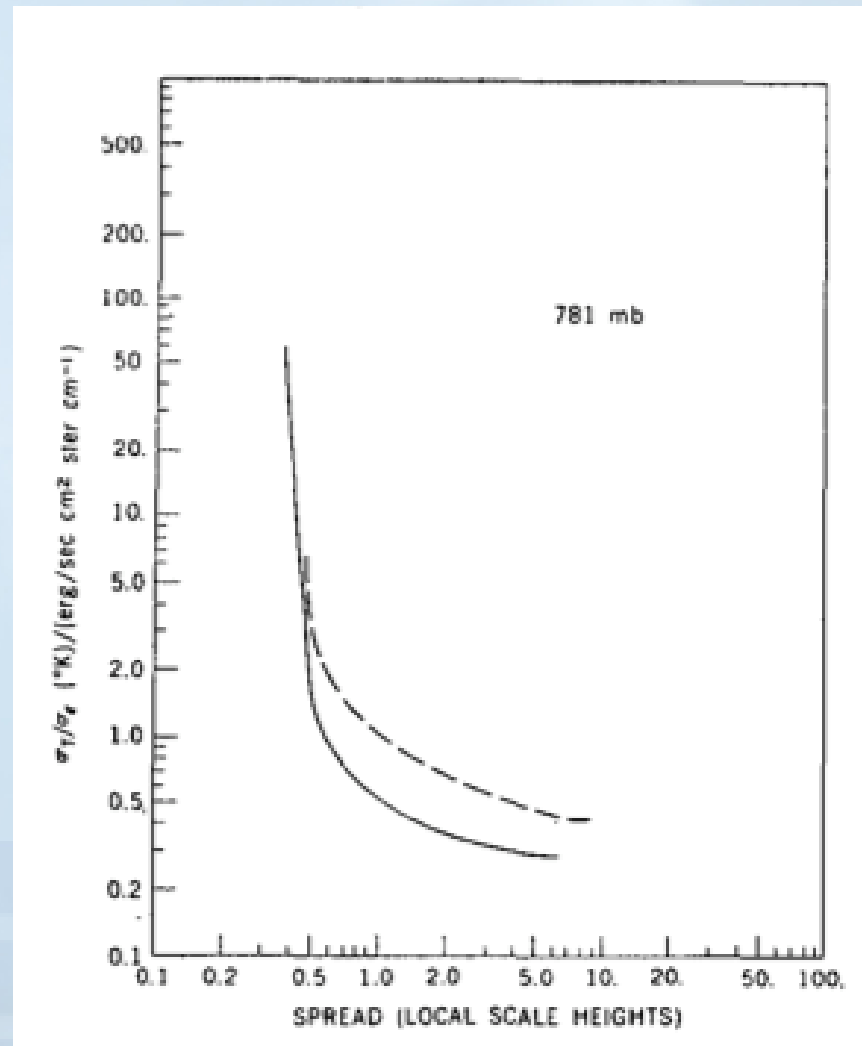
Radiometer retrieval of temperature, RH, and liquid water content during a supercooled fog event (Knupp et al., JAOT 2009)



NCAR



Tradeoff curve: Uncertainty vs Spatial Resolution



From Conrath, B., J., 1972, JAS

Summary



- Radiometer is capable of providing high temporal estimates of water vapor.
- Inherent spatial or vertical resolution of water vapor estimate is limited.
- Vertical resolution of the water vapor can be improved by the following:
 - (i) narrow-beam radiometer
 - (ii) 1-D Var method
 - (iii) including scanning measurements to vertical pointing observations
 - (iv) tradeoff between std. error in WV and spatial resolution

Summary of spatial and temporal resolutions



NCAR

3-D WATER VAPOR MEASUREMENT TECHNIQUES

Sensor	Horizontal Resolution (km)	Vertical Resolution (km)	Temporal Resolution (hr)	Frequency Band
GPS Ground Network	50	0.5 -1 Expected	0.5	L-band
Radiosondes	~315 km spacing	0.1-0.5	12	N/A
COSMIC	200-600	0.1 – 0.5	0.5 (2-hr lag)	L-band
AMSU-B	20	2	12	G-band (183 GHz)
Network of CMR-Hs	0.5	0.5-1	0.16-0.25	K-band

Radiation Transfer Equation



Radiation transfer equation:

$$\frac{dI_\nu}{ds} = -I_\nu\alpha + S$$

Solution to the above:

$$I_\nu(0) = I_\nu(s_0)e^{-\tau(s_0)} + \int_0^{s_0} B_\nu(T)e^{-\tau(s)}\alpha ds$$

Rayleigh-Jeans limit to Plank's function:

$$B_\nu(T) = \frac{2h\nu^3}{c^2} \frac{1}{e^{\frac{h\nu}{kT}} - 1} \approx \frac{2h\nu^3}{c^2} \frac{kT}{h\nu} = \frac{2kT}{\lambda^2}$$

Microwave radiometry:

$$T_b(\theta) = T_{bg}e^{-\tau(\theta)} + T_m(1 - e^{-\tau(\theta)})$$



Mathematical model for describing spatial resolution

Measurement:

.... (1)

Retrieval of WV:

.... (2)

.... (3)

..... (4)